

# Hydrocode modeling of ice-penetrating impacts on Jupiter's moon, Europa

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## Methods

We used the iSALE hydrocode (Amsden et al. 1980, Ivanov et al. 1997, Wünnemann et al. 2006) to produce full-scale models of impacts at Europa. We used a cometary impactor with density  $910 \text{ kg/m}^3$ , and physical parameters as outlined in Table 2. We validated our model by reproducing the depth-diameter ratios of mapped craters on Europa (Schenk 2002), and then ran experiments in which we tested for ice response and impact site morphology under different sets of impact conditions (Table 1). To reproduce the multi-ringed basin Tyre, we used a bolide size of 2 km (calculated based on Tyre's crater diameter, using equations in Zahnle et al. 2003), and a crust thickness of 20 km (based on Schenk's (2002) analysis of crust thicknesses beneath European craters). We used the nominal average impact velocity at Europa,  $26.5 \text{ km/s}$  (Zahnle et al. 2003).

## Introduction

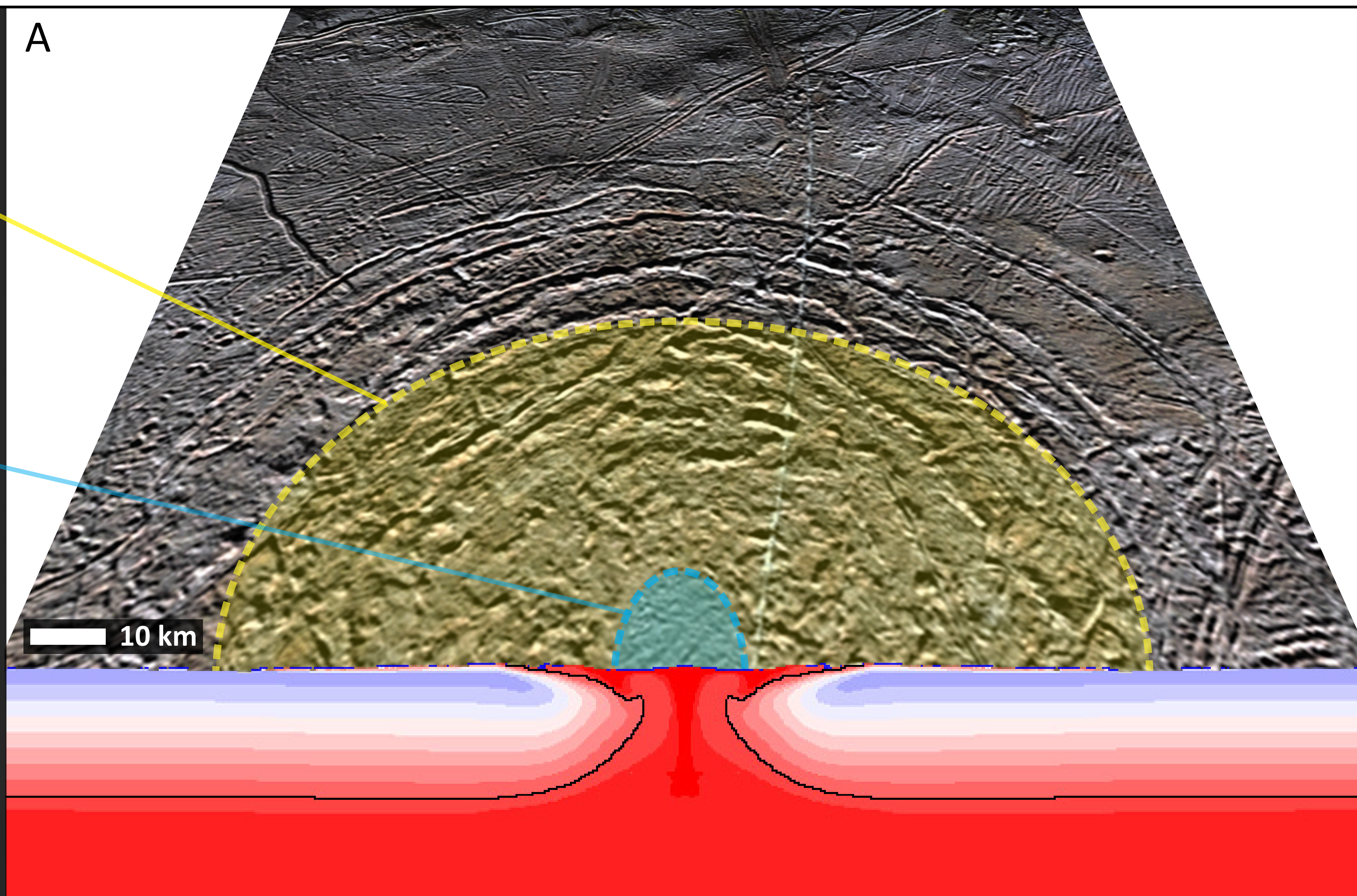
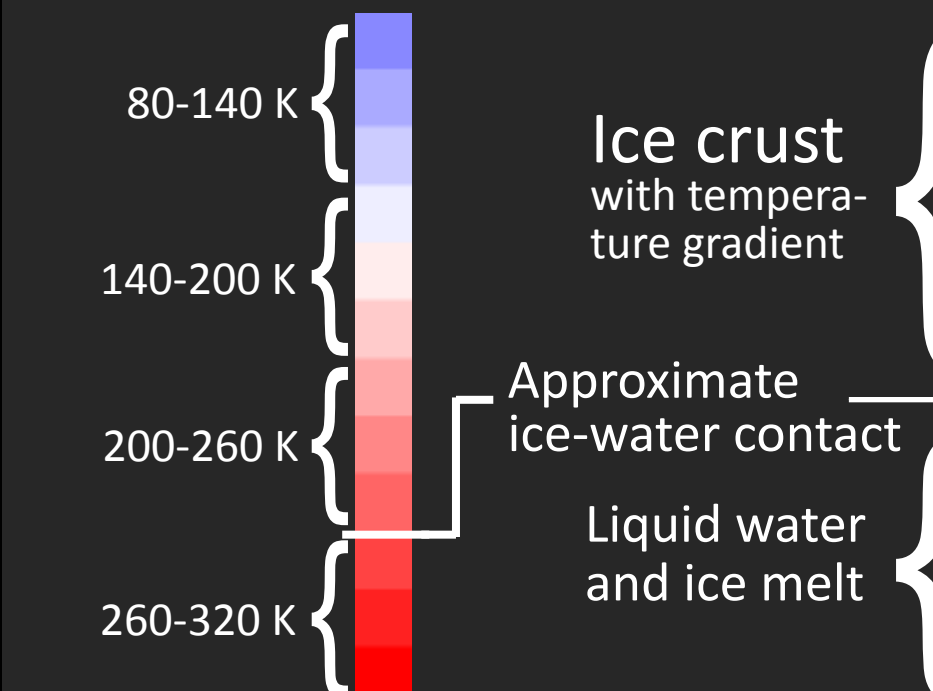
Europa's ice crust appears to overlie a water ocean (Schubert et al. 2009). If ice thickness is  $\leq 20 \text{ km}$  (indicated by crater depth-diameter ratios and lithostatic stress constraints: Schenk 2002), impacts of sufficient energy could break through to the underlying water; but little is known about the impact conditions under which this could occur, or what the morphology of resulting impact sites would be. We have modeled combinations of bolide sizes and crust thicknesses to create a matrix of impact outcomes (Table 1). We have also investigated the nature of Europa's multi-ringed basins, focusing on Tyre, and our results suggest that the impact event caused complete melt-through of the underlying crust.

## Modeling accurately reproduces Europa's Tyre

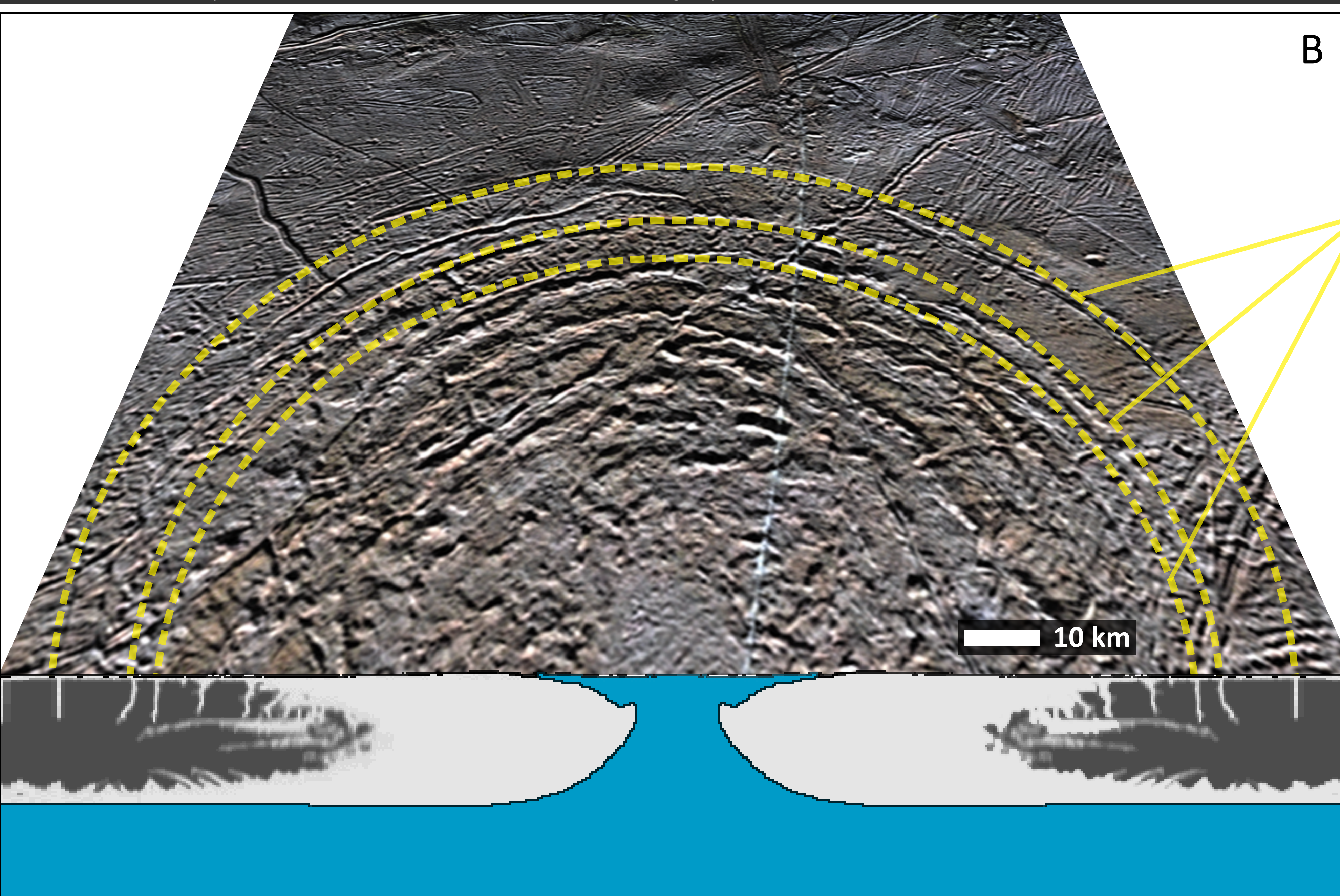
OUR MODEL MATCHES THE IMPACT FEATURES VISIBLE IN GALILEO IMAGERY OF TYRE.

Disrupted terrain with larger topography corresponds to the modeled elevation and out-flow.

The more uniform terrain at the basin center is located over the melt-through column.



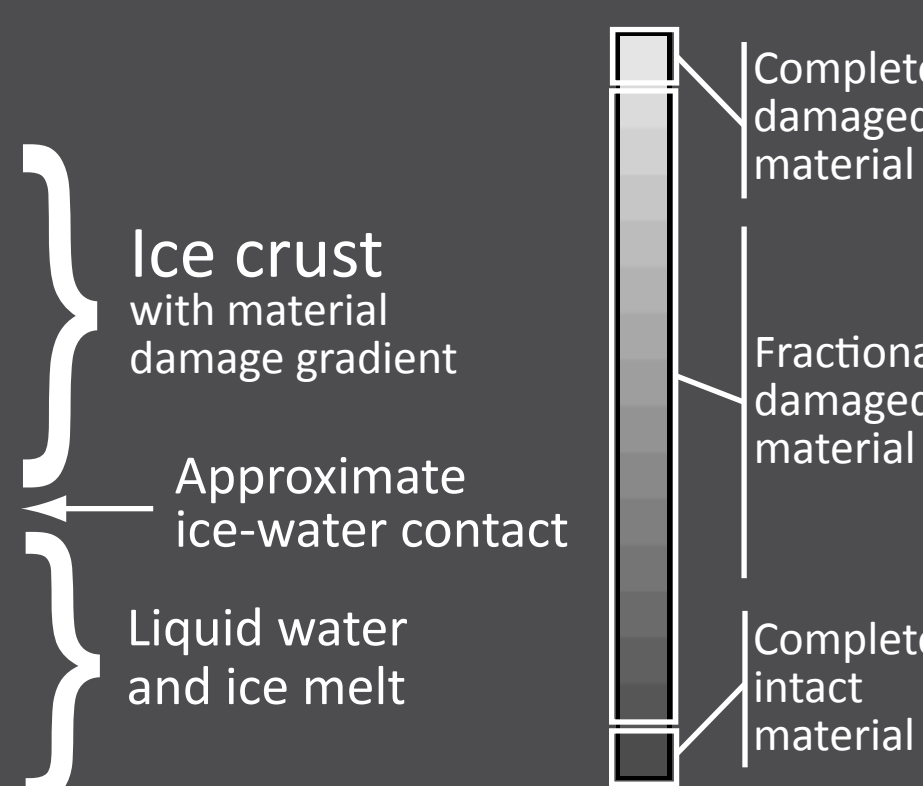
The block diagrams A and B show our model output 1500 seconds after impact along with Galileo images of Tyre at the same scale. A is color-coded for temperature and B is shaded to show the degree of crustal damage. The model (a 2 km diameter comet impacting 20 km ice over water, with impact velocity  $26.5 \text{ km/s}$ ) indicates that impact melting extends through the full thickness of the crust, that meltwater outflows beyond the crater rim, and that concentric fractures are created 70-90 km from the crater's center. The locations of crater rim and ring fractures in our models correspond well with those in the Galileo imagery.



OUR MODEL CREATES FRACTURES SIMILAR TO THOSE NEAR TYRE.

Concentric fractures surrounding Tyre correspond to fractures visible as damaged material in the model output.

Damage is a material attribute used to modify the yield strength at a point.



## Impact site morphologies

Crust Thickness (km)	Impactor Diameter (meters)							
	50	500	800	1000	1500	2000	3000	5000
1		P		P				
2		P						
3		P						
4		P						
5	C	M	P	P		P		
10	C	C	M	M	P	P		
15	C		C	C	M	P		
20	C		C	C	C	M	P	P
30						C		

**Table 1.** A summary of the outcomes of our simulations classified as either crater-forming (C), melt-through (M), or penetrating (P).

Variable	Description	Value
$\nu$	Poisson ratio	0.33
$Y_0$	Strength of intact ice at zero pressure (Senft 2009)	10 MPa
$Y_c$	Strength of damaged ice at zero pressure (Senft 2009)	0.0 MPa
$Y_m$	Strength of intact ice at infinite pressure (Senft 2009)	115 MPa
$\mu_i$	Internal friction coefficient for intact ice (Senft 2009)	2.0
$\mu_d$	Internal friction coefficient for damaged ice (Senft 2009)	0.55
$T_m$	Melting temperature	273 K
$\xi$	Thermal softening parameter	1.2
$k$	Thermal conductivity (Barr and Showman 2009)	$3.3 \text{ W/m/k}$
$\gamma_n$	Scaling parameter for A.F. viscosity (Wünnemann and Ivanov 2003)	0.1
$\gamma_T$	Scaling parameter for A.F. decay constant (Wünnemann and Ivanov 2003)	150
$C_{vib}$	Maximum vibration particle velocity (Senft 2009)	0.25
$T_{off}$	Time after which no new A.F. pressure vibrations are generated (Senft 2009)	60 s

**Table 2.** The material parameters our models used to simulate the ice crust and bolide. These include strength, temperature and acoustic fluidization (A.F.) parameters.

### Take home message:

Our numerical models accurately reproduce the geometry of Tyre, and therefore suggest that full-crustal melt-through occurred during formation of the structure. The 2-km bolide represented by our Tyre model has a return interval of  $<10 \text{ m.y.}$ , so larger impacts have definitely occurred on Europa. Our models of a range of impacts indicate that these could produce impact breaches of the crust, which would result in open holes in the ice, possibly represented by chaos terrain.

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